

1 **Tectonostratigraphy and the petroleum systems in the Northern sector of the**  
2 **North Falkland Basin, South Atlantic.**

3  
4 **Darren J. R. Jones**<sup>1\*</sup>, Dave J. McCarthy<sup>1</sup>, Thomas J.H. Dodd<sup>1</sup>

5 <sup>1</sup>British Geological Survey, the Lyell Centre, Research Avenue South, Edinburgh, EH14  
6 4AP, United Kingdom, [darjones@bgs.ac.uk](mailto:darjones@bgs.ac.uk)

7  
8 **Keywords**

9 Tectonostratigraphy, petroleum systems, North Falkland Basin, South Atlantic

10  
11 **Abstract**

12 The North Falkland Basin represents one of the frontier areas for hydrocarbon exploration in  
13 the South Atlantic. This study presents the results of new subsurface mapping using 2D  
14 seismic data in the north of the Falkland Islands offshore area, which has delineated a series  
15 of discrete grabens northwards of the main North Falkland Basin, referred collectively to as  
16 the Northern sector of the North Falkland Basin (NNFB). Six regionally significant seismic  
17 reflectors are interpreted within this data, dividing the sedimentary fill into six  
18 tectonostratigraphic packages, including: early syn-rift; late syn-rift; transitional unit; early  
19 post-rift; middle to late post-rift; and a sag unit. Structural interpretation of the 2D seismic  
20 data has led to the definition of four north-south orientated depocentres, namely: (1) the  
21 Eastern Graben, largest of the depocentres; 20km wide by 45km long, reaching depths of  
22 3km; (2) the Eastern Graben Splay, a smaller depocentre; 10km wide by 20km long,  
23 reaching depths of 2-2.5km; (3) the Western Graben Splay, the smallest depocentre; 5km in  
24 width and 20km long, with a basin depth of 2km and (4) the newly defined Phyllis Graben,  
25 which is 13km wide and 30km long, with a basin depth of 3km. A network of NW-SE and NE-  
26 SW trending faults controls the development of these grabens, separated by a Western,  
27 Eastern and Intra-Basin high. These grabens represent a northern continuation of the  
28 Northern Falkland Basin to the south. Hydrocarbon discoveries to the south of this study  
29 area (e.g. Sea Lion, Casper, Beverley, Zebedee, Isobel Deep, and Liz) confirm a working  
30 petroleum system adjacent to the Northern sector. This study has identified a number of

31 seismic anomalies, including amplitude brightening events, which potentially correspond to  
32 an extension of this petroleum system, indicating active migration pathways. The main  
33 targets, in terms of hydrocarbon interest in the northern sector, are likely to be  
34 stratigraphically trapped hydrocarbon accumulations, contained within vertically-  
35 amalgamated turbidite fan sandstone reservoirs, deposited within the early post-rift. A  
36 second, yet to be tested, syn-rift play, in which the trapping geometries are structural and the  
37 reservoirs are fluvial sandstones is also identified.

## 38 **1. Introduction**

39 The Falkland Islands offshore designated area for exploration covers approximately  
40 ~460,000 km<sup>2</sup> and has received relatively little attention in terms of hydrocarbon exploration.  
41 It is composed of four main sedimentary basins of Mesozoic to Cenozoic-age; namely the  
42 North Falkland, Falkland Plateau, South Falkland and Malvinas basins, which lie north, east,  
43 south and south-west of the islands respectively (Fig. 1). The most extensively explored and  
44 so far successful of these basins in terms of hydrocarbon prospectivity is the North Falkland  
45 Basin (NFB). More specifically, the Eastern Graben of the NFB, (Fig. 2) which has been the  
46 main focus of hydrocarbon exploration since the 1990s (Lohr and Underhill, 2015; MacAulay,  
47 2015; Richards and Hillier, 2000a and b; Thomson and Underhill, 1999; Williams and  
48 Newbould, 2015). Commercial interest, in terms of hydrocarbon potential of the NFB, has  
49 grown considerably with a number of successful exploration campaigns between 2010-2015.

50 Initial exploration of the NFB between 1998 and 1999 focused on targeting late post-rift  
51 sandstones draped over structural highs in the central parts of the NFB (MacAulay, 2015),  
52 an exploration strategy influenced by North Sea-style tilted fault block plays (Richards et al.  
53 1996a). Despite encountering an excellent, organic-rich, Lower Cretaceous-aged lacustrine  
54 source rock (up to 7.5% TOC) during drilling (Richards and Hillier, 2000b; Farrimond et al.,  
55 2015), this campaign did not encounter economical resources of hydrocarbons. However,  
56 the presence of oil and gas shows in several wells indicated a number of elements of a  
57 working petroleum system, including a mature source rock; reservoir potential sandstones

58 and a competent seal. The quantity of oil expelled from the source rock into the NFB is  
59 estimated to be approximately 60 billion barrels of oil (Richards and Hillier, 2000b).

60 Subsequently, exploration concepts shifted to basin margin-derived early post-rift  
61 sandstones. In particular, the reservoir concepts of Richards et al. (2006) described basin-  
62 margin attached fans prograding into lacustrine waters, ranging from alluvial fan, fan delta to  
63 deep-lacustrine fan systems, forming at various palaeo-water depths (Richards et al., 2006).  
64 Seismically bright amplitude anomalies, identified on 3D seismic data were, described by  
65 Richards et al. (2006), indicated various potential sediment entry points.

66 The 2010-2011 exploration campaign, was successful in discovering commercial  
67 quantities of hydrocarbons in the NFB and proved the basin margin-derived reservoir  
68 concept (MacAulay, 2015; Richards et al., 2006). This campaign targeted easterly-derived  
69 turbidite fan deposits (Bunt, 2015; Williams, 2015; Dodd et al., 2019), which form a stacked,  
70 margin-fringing succession within the Lower Cretaceous packages of the early post-rift (Fig.  
71 3), along the Eastern Flank of the North Falkland Basin's Eastern Graben (Fig. 2). The major  
72 success of the Sea Lion discovery (Francis et al., 2015; Griffiths, 2015; MacAulay, 2015;  
73 Williams, 2015) was a turning point for exploration success within the basin. Following Sea  
74 Lion , a number of analogous targets were drilled between 2010-2011 within the same play,  
75 leading to the discovery of hydrocarbons within the Casper and Beverley fans (Bunt, 2015).  
76 More recently, three wells were drilled in the NFB, leading to further discoveries in the early  
77 post-rift, such as the Zebedee and Isobel Deep Fans in 2015. These discoveries not only  
78 extended the spatial and stratigraphic extent of the petroleum system, they highlighted the  
79 further potential for future significant discoveries in the North Falkland Basin.

80 One area that has remained underexplored since the initial campaign in 1998 is the  
81 NNFB (Fig. 2), which is essentially an extension of the main NFB, and likely contains a  
82 succession of early post-rift lacustrine sediments, similar to those mapped in the Eastern  
83 Graben of the NFB to the south (Fig. 2). The stratigraphy of the NNFB also contains a

84 presumably older, syn-rift succession, which is structurally complex and remains completely  
85 un-explored.

86 This study addresses the following key questions:

- 87 1. What is the structural configuration of the NNFB?
- 88 2. What are the main controls on the structural configuration (i.e. timing and style of  
89 faulting)?
- 90 3. How does the basin configuration and fill compare and contrast with the Eastern  
91 Graben towards the south?
- 92 4. What is the nature of the tectonostratigraphy of the grabens in this area?
- 93 5. What are the likely petroleum systems and plays in the NNFB?

## 94 **2. Geological Background**

95 The NFB, described as a failed rift system (Richards et al., 1996a and b; Richards and  
96 Fannin, 1997; Lohr and Underhill, 2015), comprises a series of depocentres following two  
97 dominant structural trends: N-S oriented faulting is predominant in the northern area; whilst  
98 significant WNW-ESE oriented faults control the Southern North Falkland Basin (Fig. 2).  
99 Initial rifting of the NFB is likely to have initiated in the late Jurassic or early Cretaceous  
100 (Richards and Fannin, 1997). This rifting phase was followed by a thermal sag phase that  
101 began in the Berriasian-Valanginian (Richards, 2002). The environment of deposition  
102 throughout this sag phase is thought to be predominantly continental and deep lacustrine  
103 until Albian-Cenomanian times, when the basin began to develop increasingly marine  
104 conditions (Richards et al., 1996a and 1996b; Richards and Fannin, 1997; Richards and  
105 Hillier, 2000a).

106 The main depocentre of the NFB is orientated N-S, is approximately 30 km wide and 250  
107 km long, referred to here as the Eastern Graben (Fig. 2). A shallower depocentre is present  
108 towards the west, termed here the Western Graben, and is separated from the main Eastern  
109 Graben by an intra-graben high known as the Orca Ridge. In this Eastern Graben, the basin



110 displays an asymmetric half-graben geometry which is downthrown to the east (Fig.3a; Dodd  
111 et al., 2019; Richards et al., 1996a and 1996b; Richards and Fannin, 1997; Richards and  
112 Hillier, 2000a; Lohr and Underhill, 2015). The footwall to main basin bounding faults are  
113 composed of a Devonian-Permian platform (Richards et al., 1996a). In addition, there are a  
114 number of subsidiary depocentres immediately east of the Eastern Graben, all of which  
115 follow a similar N-S trend (Figs. 1 and 2).

116 The Southern North Falkland Basin (SNFB) represents an area intersected by a series of  
117 en-echelon WNW-ESE faults, which are easily identifiable on seismic data and gravity data  
118 (Fig. 2). The WNW-ESE faults are typically offset by the main N-S faults, suggesting two  
119 significant and distinct phases of extension, potentially associated with separate phases of  
120 rifting (Bransden et al., 1999). The older, WNW-ESE faults are similar in orientation to the  
121 trend of the Palaeozoic thrust sheets developed to the NW of the islands (Richards and  
122 Fannin, 1997; Storey et al., 1999). The WNW-ESE faults were possibly formed by  
123 reactivation of the onshore structures (Richards et al., 1996; Aldiss and Edwards, 1999).  
124 Although no well data exists in this part of the basin, the timing of this reactivation and basin  
125 development is thought to be coeval with the initial development of South Africa's Outeniqua  
126 basin during the Kimmeridgian (Thomson, 1998; Broad et al., 2012; Stanca et al., 2019).

### 127 *2.1 Seismic stratigraphy of the North Falkland Basin*

128 A tectonostratigraphic model for the NFB was presented by Richards and Hillier (2000a).  
129 The eight tectonostratigraphic units identified are: pre-rift/basement; early syn-rift; late syn-  
130 rift; transitional/sag; early post-rift; middle post-rift; late post-rift; and a post uplift sag phase  
131 (Fig. 4; Richards and Hillier, 2000a). The post rift succession is further divided into a number  
132 of sub-units, including: LC2, LC3 and LC4 in the early post-rift; LC5, LC6 and LC7 in the  
133 middle post-rift; and L/UC1 and UC1 in the late post-rift, where 'LC' is Lower Cretaceous and  
134 'UC' is Upper Cretaceous (Fig. 4). Previous seismic interpretation studies of the NFB have  
135 discussed different stratigraphic schemes (Fig. 4; Lawrence et al., 1999; Lohr and Underhill,

136 2015; Lorenzo and Mutter, 1988; MacAulay, 2015; Richards et al., 1996; Richards and  
137 Fannin, 1997 and Richardson and Underhill, 2002).

138 The pre-rift (basement) has been encountered in one well in the basin (14/09-1) which  
139 targeted an intra-basin high. At this well location (see Fig.2), the pre-rift comprises Devonian  
140 to Jurassic lithologies (Richards and Hillier, 2000a), and as a consequence of limited data,  
141 remains poorly understood.

142 The earliest phase of rifting initiated in the late Jurassic, and lasted until the Lower  
143 Cretaceous (Richards, 2002). During this time, an early to late syn-rift succession,  
144 comprising conglomerates, sandstones, organic-rich mudstones, and reworked tuffs, was  
145 deposited in a fluvial to lacustrine environment (Richards and Hillier, 2000a). Subsequently,  
146 the basin experienced a transitional-sag phase in which a succession of organic-rich  
147 lacustrine claystones were deposited (Richards et al., 1996a).

148 A succession of early post-rift sediments was deposited during the early Cretaceous  
149 (Berriasian to Aptian), resulting in a laterally and vertically extensive lacustrine mudstone  
150 and sandstone succession (Richards and Hillier, 2000a). Sediments were transported into  
151 the basin through fluvial-deltaic systems prograding from the northern-most extent of the  
152 basin, along the Eastern Graben axis, towards the south. Concomitant with this, sands were  
153 also transported into the Eastern Graben from the flanks, along feeder systems that fed a  
154 series of turbidite fans; creating a complex, heterogeneous succession of sandstones and  
155 interbedded mudstone facies (as described in Dodd et al., 2019). In particular, the easterly-  
156 derived sandstones currently represent the main reservoir lithologies identified in the NFB, to  
157 date.

158 During the Lower Cretaceous (Aptian-Albian) the basin fill began to develop as a thick  
159 succession of middle to late post-rift sediments were deposited. Here the sediment fill is  
160 characterised by a transition from a lacustrine dominated succession to terrestrial-fluvial  
161 systems (Richards and Hillier, 2000a).

162 In the late post-rift (Albian to Palaeocene) the basin experienced the first significant  
163 marine transgression. The resulting sediments comprised of claystone interbedded with  
164 sandstone, deposited in a restricted, marginal marine or lagoonal environment (Richards and  
165 Hillier, 2000a). Following the late post-rift (Palaeocene) the region underwent significant  
166 uplift, during which up to 800 m of overburden is thought to have been removed from parts of  
167 the basin (Richards et al., 1996a and b). The post uplift sediments consist of a dominant  
168 succession of claystone with interbedded sandstone, deposited in a fully developed marine  
169 basin environment (Richard and Hillier, 2000a).

### 170 **3. Datasets and Methodology**

171 This study uses 1,250 km of 2D seismic reflection data ( the “FALK2000” survey)  
172 collected and processed by Veritas in 2000 on behalf of Lasmco plc, located north of  
173 operated blocks PL001, PL032 and PL033 (Fig. 5). The seismic data is post stack time  
174 migrated, which displays a positive polarity and the data is zero-phased. These seismic lines  
175 have a line spacing of 2.5-5 km in an N-S orientation and 2.5-10 km in an E-W orientation.  
176 Overall, the quality of the 2D seismic data is of reasonable quality down to 3-3.5 seconds  
177 two-way-travel-time (TWTT). Beyond this, the signal to noise ratio increases significantly and  
178 the reflections become chaotic (Figs. 3, 6 and 7). In addition to the seismic reflection data,  
179 major structures and basins were identified using Bouguer gravity data (Fig. 2) from global  
180 marine data (Sandwell et al., 2014).

181 To date, there are no wells within the study area (Fig. 5); however, a seismic correlation  
182 has been made from the “FALK2000” 2D survey, southwards into the “Company Composite”  
183 3D seismic survey (Figs. 2 and 8) which consists of several merged 3D seismic datasets  
184 acquired by Shell in 1997; Desire in 2004 and Rockhopper in 2007 and 2011. This profile  
185 intersects the nearest well to the study area (14/05-1A) and wells near the Sea Lion  
186 discovery (14/10-2, 14/10-3, 14/10-5 and 14/10-7). In these more southerly areas, geological  
187 understanding is more mature and the stratigraphy is better constrained. These tie-lines

188 enabled seismic well picks to be interpreted across into the study area, providing some  
189 stratigraphic control on the interpretation.

190 Seismic data were interpreted using seismic and stratigraphic concepts (sensu Mitchum  
191 et al., 1977; Vail et al., 1977; Hubbard et al., 1985). TWTT surface maps were produced  
192 from the seismic interpretation, gridded at 100 m increments. Fault polygons were created  
193 through extrapolation of 2D fault segments using a standard triangulation gridding algorithm  
194 method.

#### 195 **4. Tectonostratigraphy**

196 In the NNFB, six regionally significant seismic reflections were identified within the 2D  
197 seismic data (Figs. 6 and 7), defining six tectonostratigraphic units (Fig. 4). These units have  
198 been defined by extrapolation of seismic data from the main Eastern Graben of the NFB  
199 (Fig. 8). The seismic reflectors defining these units are: top basement (TB); top early syn-rift  
200 (TESR); top late syn-rift (TLRS); top transitional/sag (TS); top early post-rift (TEPR); and top  
201 late post-rift (TLPR).

##### 202 *4.1 Basement and syn-rift (TB, TESR and TLRS)*

203 The deepest reflector that can be mapped on a regional scale, the top basement (TB),  
204 forms an unconformable surface that is present across the entire seismic survey (Figs. 6 and  
205 7). In the northern sector of the NFB, top basement is less clearly imaged below 3 seconds  
206 TWTT, while further south in the main Eastern Graben, the basin deepens with the "TB"  
207 reflector found around 4 seconds TWTT (Fig. 4). It often presents as a very bright amplitude  
208 on basement highs such as the Eastern Flank, whilst in deeper parts of the seismic data,  
209 there are small (<1 km wide), discontinuous, high amplitude reflections beneath the "TB"  
210 reflector (Fig. 6). It is possible these features could represent either igneous intrusive bodies  
211 (dykes or sills) or Devonian-Carboniferous metasediments, deposited prior to the basin rifting  
212 event (Fig. 6). Two dyke swarms have been identified onshore Falkland Islands (Stone et al.,  
213 2008; Richards et al., 2013), with ages of 188-178 Ma and 135-121 Ma.

214 The seismic reflector marking the top of the early syn-rift (TESR) is challenging to  
215 distinguish laterally. Internally, the early syn-rift often displays, high amplitude reflectors,  
216 which are divergent and mound-like in appearance (Figs. 6 and 8). In some places this unit  
217 thins onto pre-rift basement highs (Fig. 7), in other cases these high amplitude reflections  
218 are discontinuous and have a chaotic appearance (Fig. 6).

219 The top late syn-rift (TLSR) is marked by an undulating, high amplitude seismic reflector,  
220 which separates the underlying, slightly transparent late syn-rift package from the overlying  
221 transitional package (Fig. 6). In places, the late syn-rift onlaps onto the underlying early syn-  
222 rift unit (Fig. 7). The internal character of the late syn-rift is relatively transparent throughout  
223 (Fig. 6), with discrete, alternating, high and low amplitude packages observed (Fig. 7b).

#### 224 *4.2 Transitional to post-rift (T/S, TEPR and TLPR)*

225 The top of the transitional/sag (T/S) reflector is marked by a high amplitude laterally  
226 continuous reflector, which onlaps against the basin margins of the Eastern Graben, as well  
227 as the Eastern Flank (Figs. 6 and 8). The transitional/sag interval is characterised by a  
228 relatively uniform sediment thickness, which only ever thins out onto the basement highs.  
229 Internally, it contains isolated, chaotic, high amplitude events (Fig. 6).

230 The top early post-rift reflector (TEPR) is a prominent, high amplitude reflector that is  
231 laterally continuous across this seismic survey and defines the top of the early post-rift unit  
232 (Fig. 7). Internally, this unit contains high amplitude, sheet-like reflectors at the base, along  
233 with clinoform-like geometries forming at the top of the package, which appear to downlap  
234 onto the sheet-like reflectors beneath (Fig. 8). These clinoforms prograde from the north  
235 towards to the south.

236 The top late post-rift seismic reflector (TLPR) is laterally continuous across the seismic  
237 survey (Figs. 3, 6, 7 and 8). The late post-rift unit, internally, consists of laterally continuous,  
238 seismically transparent intervals at the base, developing into alternating, high and low  
239 amplitude reflectors towards the top.

240 The seismic package above the TLPR represents the post sag uplift sequence, which  
241 continues to the seabed. The package is generally transparent containing sub-parallel to  
242 parallel, low amplitude reflectors. Although within the package there are a few high  
243 amplitude, laterally continuous reflectors, which represent unconformable surfaces within the  
244 succession. This can be shown by downlap terminations of divergent reflectors on to these  
245 surfaces (Fig.6).

## 246 **5. Structural interpretation**

247 A number of two-way-travel time structural maps were produced from the interpretation of  
248 the 2D seismic data in order to understand the structural evolution of the NNFB (Fig. 9). A  
249 map of top basement (Fig. 9a) shows four N-S orientated structural lows, defined from west  
250 to east as the: Western Graben Splay; Eastern Graben; Eastern Graben Splay; and the  
251 Phyllis Graben. The Western Graben Splay, Eastern Graben and Eastern Graben Splay,  
252 together, form the northern continuation of the main graben of the NFB (Fig.2). The Western  
253 High is considered to be a northward extension of the Orca Ridge to the south and therefore  
254 the Western Graben Splay is likely to be a northward extension of the Western Graben of  
255 the NFB (Fig. 2). The Western High separates the Western Graben Splay from the Eastern  
256 Graben. The Eastern Flank forms the main structural high in the eastern part of this area.  
257 This has been separated here into a spur termed the Intra-Basin High, which separates the  
258 Eastern Graben and Eastern Graben Splay (Fig. 9a).

259 The Eastern Flank forms the main structural high on the eastern side of the Eastern  
260 Graben and continues southwards to the Sea Lion discovery area (Fig. 2). In the southern  
261 part of the NNFB, both the Eastern Graben and Eastern Graben Splay have half-graben  
262 geometries, and deepen towards the east against the main bounding faults (Figs. 8 and 9). A  
263 series of NW-SE and NE-SW orientated faults are present across the NNFB and define the  
264 structural orientation of these grabens (Fig. 9a-9e).

265 The Phyllis Graben (PG), located directly to the east of the Eastern Flank (Figs. 2 and 9),  
266 is composed of a series of half-grabens that are predominantly orientated N-S. This graben  
267 also displays an asymmetrical profile, deepening towards the north-west (Figs. 6 and 9a). It  
268 is possible that the Phyllis Graben continues north of the study area, developing into a  
269 geographically larger suite of grabens, shown as N-S oriented 'gravity-lows' in Bouguer  
270 gravity data (Fig. 2). These grabens have a comparable gravity signature to that of the  
271 Eastern Graben.

272 By the end of the early syn-rift, all four of the main structural lows had developed (Fig.  
273 9b). At this stage, the Eastern Graben was the deepest and spatially largest of the four  
274 depocentres. In the Eastern Graben Splay, the early syn-rift interval deepens towards the  
275 east and south against the Eastern Flank, while the Western Graben Splay deepens to the  
276 south (Fig. 9b). In contrast, the early syn-rift of the Phyllis Graben deepens north westerly. In  
277 general, the early syn-rift interval maintains a relatively consistent thickness in the study area  
278 (Fig. 7), but thickens southwards towards the Sea Lion Discovery (Fig. 8). In some areas  
279 (Fig. 7), the early syn-rift shallows up against the main bounding faults, particularly in the  
280 northern part of the Eastern Graben and Eastern Graben Splay.

281 The late syn-rift interval follows a similar structural pattern as the underlying early syn-rift,  
282 with increasing deepening in the centre and southern of the Eastern Graben in the NNFB. In  
283 the southern part of the NNFB, the late syn-rift onlaps the underlying early syn-rift interval  
284 against the Intra-Basin High (Fig. 7).

285 By the end of the transitional/sag phase, the sedimentary cover was significantly more  
286 extensive, with the overstepping of the Intra-Basin High between the Eastern Graben and  
287 Eastern Graben Splay and partially over the Western High (Fig. 9d). Structural depth  
288 increases southwards and towards the centre in the Eastern Graben, Eastern Graben Splay,  
289 whilst in the Phyllis Graben the basin depth increases northwards (Fig. 9d). Fault trends  
290 remain consistent with NW-SE and NE-SW trends observed at the top of the late syn-rift. A



291 network of fault terraces is present at the southern extent between the Eastern Graben and  
292 Eastern Graben Splay (Fig. 9d).

293 By the early post-rift, the basin continued to fill with sediments, primarily within the  
294 Eastern Graben and the Eastern Graben Splay. The Western Graben Splay displays a  
295 similar amount of deepening to that exhibited during the transitional/sag phase, whilst the  
296 Phyllis Graben has experienced overall deepening. During the early post-rift units, the Phyllis  
297 Graben appears structurally deeper than the Eastern Graben, the Western Graben and the  
298 Eastern Graben Splay. Furthermore, sediments have encroached further northwards onto  
299 the Intra-Basin High (Fig. 9e). The Western Graben Splay, Eastern Graben, Eastern Graben  
300 Splay and Phyllis Graben fault trends remain consistent, with a NW-SE and NE-SW trend as  
301 seen during the transitional/sag phase.

302 In the late post-rift, sediment cover is preserved across the Western and Eastern Flanks,  
303 as well as the Intra-Basin High (Fig. 9f). Here, the late post-rift sediments deepen  
304 northwards in the Eastern Graben and Phyllis Graben respectively (Fig. 9f). By this phase,  
305 most of the faulting had terminated with only a few NW-SE faults remaining active.

## 306 **6. Discussion**

### 307 *6.1 Basin development in the NNFB*

308 This study has shown that faults that were active during the syn-rift phase remained  
309 consistently active until the early post-rift phase. The NW-SE faults observed in the NNFB  
310 may represent similar structures as those observed in the Southern North Falkland Basin,  
311 which were interpreted as reactivated thrust faults similar to those seen onshore (Richards et  
312 al., 1996; Aldiss and Edwards, 1999; McCarthy et al., 2017). The NE-SW orientated faults  
313 are likely to have formed due to the initial E-W extension associated with the opening of the  
314 South Atlantic during the late Jurassic- early Cretaceous (Richards and Fannin, 1997).  
315 These faults form a component of the fault architecture along with the NW-SE faults, defining

316 the margins of the N-S trending depocentres, namely the Eastern Graben, Eastern Graben  
317 Splay, Western Graben and Phyllis Graben.

318 It is likely the initial rifting of the NNFB occurred contemporaneously with the central part  
319 of the NFB to the south. During this initial rifting, an early syn-rift phase led to the  
320 development of accommodation space within the centre of each of these grabens. This  
321 rifting continued into the late syn-rift with accommodation space increasing, particularly  
322 within the Eastern Graben (Fig. 9c). The presence of structural highs such as the Western  
323 Flank, Western High, Intra-Basin High and Eastern Flank, as well as consistent fault trends  
324 in the early and late syn-rift (Fig. 9b-9c), suggest these faults remained tectonically active  
325 throughout the syn-rift.

326 As rifting halted, and the NNFB entered the transitional/sag phase, the Western High and  
327 Western Flank became inactive and an overstepping succession was deposited across  
328 these highs. It is likely that during this time the amount of accommodation space developed  
329 at the edge of the Intra-Basin High started to be outpaced by sediment input, evidenced by  
330 the partial flooding and deposition of sediments over the high during this time. However, the  
331 Eastern Flank continued to remain a topographical high at the stage (Fig. 9d). The  
332 consistent presence of the NW-SE and NE-SW fault trends illustrates these faults remained  
333 active throughout the syn-rift into the transitional/sag phase. During the early post-rift, the  
334 Eastern Graben and Eastern Graben Splay formed a single connected depocentre  
335 deepening southwards and remained isolated from the Phyllis Graben by the Eastern Flank  
336 (Fig. 9e). Clinofolds observed within the early post-rift of the Eastern Graben (Fig. 8)  
337 suggest a prograding deltaic system drained into the basin from the north.

338 The Phyllis Graben, seems to have developed at a steady rate throughout the syn-rift  
339 and transitional/sag phase, as seen by the gradual structural deepening of the basin  
340 northwards (Fig. 9b-9d). During the early post-rift the Phyllis Graben appears to have more  
341 subsidence than the Eastern Graben (Fig. 9e), while in the late post-rift both depocentres  
342 had a consistent depth (Fig. 9f). By the late post-rift, most of this tectonic activity had ceased

343 with only a few NW-SE faults remaining active, having either exploited crustal weaknesses  
344 derived from mid post-rift faults or through differential compaction. At this stage, the Eastern  
345 Flank was covered with sediment, as the Eastern Graben and Phyllis Graben amalgamated  
346 into a single, large depocentre (Fig. 9f).

## 347 *6.2 Hydrocarbon prospectivity of the NNFB*

### 348 *6.2.1 Source Rock*

349 No well data is available in the NNFB and consequently source rock intervals to the south  
350 have been used to provide analogous data. In the North Falkland Basin, the main source  
351 rock intervals are organic rich claystones within the transitional/sag and early post-rift  
352 tectonostratigraphic units (Richards and Hillier, 2000b). These claystones were deposited in  
353 an anoxic, lacustrine environment during the Berriasian to Aptian (Richards and Hillier,  
354 2000b), and are thought to be responsible for charging the reservoirs of the Sea Lion  
355 discovery (MacAulay, 2015; Farrimond et al., 2015). The source rocks in the Early Post Rift  
356 comprise Type I and II kerogens, and generally increase in total organic carbon (TOC) from  
357 the transitional/sag unit, into the overlying early-post rift (Richards and Hillier, 2000b). Basin  
358 modelling has suggested the main phase of oil generation of the early post-rift source rock  
359 took place during the late Cretaceous between 70-100Ma (Richards and Hillier, 2000b).

360 Analysis of the 2010-2011 wells characterised the recovered oil samples as: “a dark,  
361 waxy, lacustrine oil with an API ranging from approximately 24-29° sourced from various oil  
362 families” (Farrimond et al., 2015). Figure 8 illustrates that the transitional and early post-rift  
363 units, which contain this main source rock interval, remain at the same depth across the  
364 main NFB and into the NNFB (between 1.7-2.3 seconds). Therefore, it is possible that  
365 hydrocarbons have been generated in this part of the basin. Continuous sub-parallel/parallel,  
366 low frequency, high amplitude reflectors in these units are likely to represent deep lacustrine  
367 organic rich claystone source rocks (Fig.8). In contrast, discontinuous reflectors are likely to  
368 represent shallow lacustrine sediments, which consist of organic lean claystone units  
369 interbedded with sandstone units.

370 In addition, a secondary source rock interval is likely within claystone-dominated units  
371 within a fluvial succession of the late syn-rift, deposited during the Tithonian to Berriasian  
372 (Richards and Hillier, 2000b). Rock-Eval pyrolysis studies, completed on wells 14/05-1A and  
373 14/10-1, suggest the presence of Type II source rocks in the late syn-rift with 4.5% average  
374 TOC (Richards and Hillier, 2000b).

375 Basin modelling suggested an oil window between 2,800–3,500 m (Richards and Hillier,  
376 2000b) across the NFB. In the central part of the NFB, the syn-rift package reaches depths  
377 >4,000 m, Vitrinite reflectance data suggests any source rock encountered here is likely to  
378 be within the gas window (Richards and Hillier, 2000b). However, as the syn-rift is shallower  
379 in the NNFB (c. 3,000 m) there is potential for it to be oil prone in this area.

#### 380 6.2.2 Reservoirs

381 Fluvial sandstones have the potential to act as reservoirs within the early syn-rift  
382 (Richards and Hillier, 2000b). These sandstones have been encountered in the nearest well  
383 14/05-1A, with several zones of net thicknesses reaching up to 40 m, with porosities ranging  
384 from 4.4–7.5% (Richards and Hillier, 2000b). Greater potential is likely in fluvial sandstones  
385 of the late syn-rift, where thicker successions have been encountered with up to 125 m of  
386 net sandstone and porosities ranging between 27.8–30.4% (Richards and Hillier, 2000b).

387 In the NFB, the best understood reservoir intervals are contained within the early post-rift  
388 unit, these sandstones were first identified as the primary reservoir target during the drilling  
389 campaign in 1998, and were later confirmed during the Sea Lion discovery in 2010 (Holmes  
390 et al., 2015; Williams, 2015). In the Sea Lion Fan, these reservoirs consist of well-sorted,  
391 fine- to medium grained, high-density turbidite sandstones deposited in a deep-lacustrine  
392 turbidite fan setting (Dodd et al., 2019). The fans are composed of overlapping lobes fed into  
393 the basin from the east. Reservoir quality within the sandstones in the Sea Lion Fan is  
394 generally good, with porosity and permeability values averaging 22% and 185 mD,  
395 respectively (Williams, 2015). On the 2D seismic data, similar geometries comparable to that

396 of the Sea Lion complex are observed (Fig. 8). The Northern Lead forms a discrete, 5-7 km  
397 long, high amplitude seismic event (Fig. 8), which was deposited near to the base of the  
398 early post-rift unit. In both examples of the Sea Lion complex and Northern Lead seismic  
399 reflectors are significantly stronger than the surrounding sediments and display a “mound”  
400 like topography, within an otherwise flat lying package of reflectors, suggesting depositional  
401 relief. Laterally, both display a reduction in seismic amplitudes towards the edge of each  
402 feature and are found at the base of southerly prograding clinoforms representing delta  
403 foresets. In addition, a series of high amplitude, sheet-like seismic reflectors can also be  
404 seen in the early post-rift, which may represent hydrocarbon-filled lacustrine turbidite  
405 sandstones (Fig.7).

### 406 6.2.3 Traps

407 The first phase of exploration drilling in 1998 focused on targeting structural, four-way dip  
408 closures, such as drilled by well 14/09-1, which drilled on the crest of a large, tilted fault  
409 block. One of the main reasons for failure was the ineffective top seal (Richards and Hillier,  
410 2000b). In the NNFB there are a number of potential two-way and three-way dip closures  
411 identified in the hanging walls of faults in the early and late syn-rift intervals (Fig. 9b and 9c),  
412 which are yet to be tested.

413 In the early post-rift, there is potential for stratigraphic traps containing deltaic-top and  
414 delta-front sandstones. These sediments are likely to be part of prograding deltaic systems,  
415 which can be observed as clinoform geometries in the seismic data (Fig. 8). Furthermore,  
416 the more distal delta deposits, which are likely to be more mud-prone, provide lateral seal  
417 potential these trapping geometries.

418 To date the most successful trapping geometries in the North Falkland Basin are  
419 complex, combined structural and stratigraphic traps, particularly within the early post-rift  
420 (e.g., Sea Lion discovery; MacAulay, 2015; Dodd et al., 2019). Firstly, the stratigraphic  
421 component is provided by deep lacustrine turbidite fans that display abrupt, lateral pinch-

422 outs and up-dip sealing through the detachment of feeder systems facilitated through slope  
423 bypass (Dodd et al., 2019). Secondly, the structural component is provided through the  
424 draping of turbidite sands along basin margin geometries and over the inversion-related high  
425 in the centre of the Eastern Graben (Fig. 3). Finally, in places basin-margin faults aid up-dip  
426 sealing through the offsetting of turbidite fan feeder channels from the depositional lobes,  
427 providing an element of fault closure to some of these traps.

#### 428 *6.2.4 Seal*

429 The regional seal across the NFB is formed by a thick mudstone succession within the  
430 early post-rift (Richards and Hillier, 2000b). The early post-rift unit is laterally extensive  
431 across the grabens of the NNFB according to the correlated seismic package (Fig. 9d),  
432 therefore likely forming a regionally effective seal. The seal will be most effective in the  
433 centre of the basin depocentres where the mudstone accumulations are likely to be thicker  
434 (Richards and Hillier, 2000b). In the NNFB there is also potential for sealing mudstones  
435 within the middle post-rift and late post-rift units (Richards and Hillier, 2000b).

#### 436 *6.2.5 Seismic evidence for an active petroleum system*

437 Evidence for an active hydrocarbon system in the NFB is proven by the by the  
438 discoveries to the south. It is inferred in the NNFB from seismic anomalies present in the  
439 seismic data. In the NNFB, the early post-rift unit is intersected by major, deep-seated  
440 normal faults that penetrate through to the late post-rift, typically at the basin margins (Fig.  
441 7). Amplitude anomalies with a negative impedance contrast, or 'soft-kicks', are common,  
442 some of which can be interpreted as bright-spots (Fig. 10a). The bright spots in the Eastern  
443 Graben of the NNFB appear to occur within the middle post-rift unit, brightening near faults  
444 (Fig. 10a), which may have acted as fluid-conduits or traps for hydrocarbon migration (Fig.  
445 10a). In addition, gas chimney features are observed cross-cutting seismic reflections (Fig.  
446 10b). Paleo-pockmarks visible in the seismic (Fig. 10b) could indicate thermogenic or  
447 biogenic gas associated with the deeper source rock intervals (Cartwright and Santamarina,

448 2015). Stratigraphic packages within the early post-rift show brightening along reflections  
449 (Fig. 10c), which may indicate fluid filled sandstone packages within a turbidite fan  
450 succession, similar to those of the Sea Lion Fan (Dodd et al., 2019).

#### 451 6.2.6 Summary of Petroleum System

452 The conceptual model for the petroleum system of the NNFB is summarised in Figure 11.  
453 In this area, the best reservoirs are likely to be within early post-rift structural-stratigraphic  
454 traps, with sand-rich turbidite fans, and the fluvial sandstones of the syn-rift in structural  
455 traps and in. The most organic rich sediments are likely to be found in the centre of the  
456 graben, in the transitional/sag unit, whilst a secondary organic rich interval maybe present  
457 along the hanging wall margins which generally represent the deepest section of the  
458 grabens during the late syn-rift unit. Moreover, both source rock intervals in the syn-rift and  
459 transitional/sag unit are likely to be mature in the NNFB as they are situated within the  
460 estimated oil window (2,250-3,000m). Finally, hydrocarbon migration potential along major  
461 faults, through and above the source/seal interval, into the overlying, thick deltaic front  
462 sandstones.

### 463 7. Conclusions

- 464 1. This study has defined the structural configuration of the Northern sector of the North  
465 Falkland Basin. The NNFB consists of two main depocentres: (a) a Western Graben  
466 and a continuation of the Eastern Graben of the NFB; and (b) a newly defined  
467 depocentre, the Phyllis Graben.
- 468 2. A series of NW-SE and NE-SW trending faults control the development of the  
469 grabens throughout the syn-rift until the late post-rift when tectonic activity ceased  
470 and sedimentation covered the entire NNFB.
- 471 3. The stratigraphy of the NNFB is separated into six tectonostratigraphic units, defined  
472 in seismic data as: early syn-rift; late syn-rift; transitional/sag; middle post-rift; late  
473 post-rift and overlying sag unit.



- 474 4. Detailed mapping has identified two hydrocarbon plays for the NNFB: i) a fluvial, syn-  
475 rift structural leads, in either 2 or 3-way dip hanging wall closures against faults; ii) an  
476 early post-rift, combined structural-stratigraphic play, which may relate to turbidite fan  
477 sandstones sourced from the north or east, analogous to well documented reservoirs  
478 found in the Sea Lion Main Complex.
- 479 5. The main source rock for the NNFB is potentially organic rich lacustrine deposits  
480 within the transitional/sag unit located in the Western Graben Splay, Eastern Graben,  
481 Eastern Graben Splay and Phyllis Graben.

482

### 483 **Acknowledgements**

484 This paper is published by permission of the Director of Mineral Resources, Falkland Islands  
485 Government, and the Executive Director, British Geological Survey (UKRI). Margaret  
486 Stewart and Dave Schofield are specifically thanked for their helpful comments. Chris Elders  
487 and Nicola Scarselli are thanked for their constructive reviews of this paper. We would also  
488 like to extend our gratitude to Tiago Alves for his editorial support.

489

### 490 **References**

- 491 Aldiss, D.T. and Edwards, E.J., 1999. The geology of the Falkland Islands. *British Geological*  
492 *Survey Technical Report WC/99110*.
- 493 Bransden, P.J.E., Burges, P., Durham, M.J. and Hall, J.G., 1999. Evidence for multi-phase  
494 rifting in the North Falklands Basin. *Geological Society, London, Special*  
495 *Publications*, 153(1), pp.425-443.
- 496 Broad, D.S., Jungslager, E.H.A., McLachlan, I.R., Roux, J. and Van der Spuy, D., 2012.  
497 South Africa's offshore Mesozoic basins. In *Regional Geology and Tectonics: Phanerozoic*  
498 *Passive Margins, Cratonic Basins and Global Tectonic Maps* (pp. 534-564).

- 499 Bunt, R. J. W. 2015. The use of seismic attributes for fan and reservoir definition in the Sea  
500 Lion Field, North Falkland Basin, *Petroleum Geoscience*, **21**, 137-149.
- 501 Cartwright, J. and Santamarina, C. 2015. Seismic characteristics of fluid escape pipes in  
502 sedimentary basins: Implications for pipe genesis. *Marine and Petroleum Geology*, **65**, 126-  
503 140.
- 504 Dodd, T.J.H., McCarthy, D.J. and Richards, P.C. 2019. A depositional model for deep-  
505 lacustrine, partially confined, turbidite fans: Early Cretaceous, North Falkland Basin.  
506 *Sedimentology*, 66 (1), pp.53-80.
- 507 Farrimond, P., Green, A. and Williams, L. 2015. Petroleum geochemistry of the Sea Lion  
508 Field, North Falkland Basin. *Petroleum Geoscience*, **21**, 125-135.
- 509 Francis, A., Lewis, M. and Booth, C., 2015. Sea Lion Field, North Falkland Basin: seismic  
510 inversion and quantitative interpretation. *Petroleum Geoscience*, pp.2014-048.
- 511 Griffiths, A., 2015. The reservoir characterization of the Sea Lion Field. *Petroleum*  
512 *Geoscience*, pp.2014-041.
- 513 Holmes, N., Atkins, D., Mahdi, S. and Ayress, M., 2015. Integrated biostratigraphy and  
514 chemical stratigraphy in the development of a reservoir-scale stratigraphic framework for the  
515 Sea Lion Field area, North Falkland Basin. *Petroleum Geoscience*, 21(2-3), pp.171-182.
- 516 Hubbard, R.J., Pape, J. and Roberts, D.G., 1985. Depositional Sequence Mapping as a  
517 Technique to Establish Tectonic and Stratigraphic Framework and Evaluate Hydrocarbon  
518 Potential on a Passive Continental Margin: Chapter 5.
- 519 Lawrence, S.R., Johnson, M., Tubb, S.R. and Marshallsea, S.J., 1999. Tectono-stratigraphic  
520 evolution of the North Falkland region. *Geological Society, London, Special*  
521 *Publications*, 153(1), pp.409-424.
- 522 Lohr, T. and Underhill, J. 2015. Role of rift transection and punctuated subsidence in the  
523 development of the North Falkland Basin. *Petroleum Geoscience*, **21**, 85-110.

- 524 Lorenzo, J.M. and Mutter, J.C. 1988. Seismic stratigraphy and tectonic evolution of the  
525 Falkland/Malvinas Plateau. *Revista Brasileira de Geociencias*, **18(2)**, 191-200.
- 526 MacAulay, F. 2015. Sea Lion Field discovery and appraisal: a turning point for the North  
527 Falkland Basin. *Petroleum Geoscience*, **21**, 111-124.
- 528 McCarthy, D., Aldiss, D., Arsenikos, S., Stone, P. and Richards, P., 2017. Comment on  
529 "Geophysical evidence for a large impact structure on the Falkland (Malvinas) Plateau".  
530 *Terra Nova*, 29(6), pp.411-415.
- 531 Mitchum Jr, R.M., Vail, P.R. and Sangree, J.B., 1977. Seismic stratigraphy and global  
532 changes of sea level: Part 6. Stratigraphic interpretation of seismic reflection patterns in  
533 depositional sequences: Section 2. Application of seismic reflection configuration to  
534 stratigraphic interpretation.
- 535 Richards, P., 2002. Overview of petroleum geology, oil exploration and associated  
536 environmental protection around the Falkland Islands. *Aquatic Conservation: Marine and*  
537 *Freshwater Ecosystems*, 12(1), pp.7-14.
- 538 Richards, P., Duncan, I., Phipps, c., Pickering, G., Grzywacz, J., Hout, R. & Meritt, J. (2006)  
539 Exploring for Fan Delta Sandstones in the Offshore Falkland Islands. *Journal of Petroleum*  
540 *Geology*, Vol. 29, pp. 199-214.
- 541 Richards, P.C. and Hillier, B.V. 2000a. Post-drilling analysis of the North Falkland Basin –  
542 Part 1: Tectono-stratigraphic framework. *Journal of Petroleum Geology*, **23(3)**, 253-272.
- 543 Richards, P.C. and Hillier, B.V. 2000b. Post-drilling analysis of the North Falkland Basin –  
544 Part 2: Petroleum system and future prospects. *Journal of Petroleum Geology*, **23(3)**, 273-  
545 292.
- 546 Richards, P.C. and Fannin, N.G.T. 1997. Geology of the North Falkland Basin. *Journal of*  
547 *Petroleum Geology*, **20(2)**, 165-183.

- 548 Richards, P.C., Gatliff, R.W., Quinn, M.F. and Fannin, N.G.T. 1996a. Petroleum potential of  
549 the Falkland Islands offshore area. *Journal of Petroleum Geology*, **19(2)**, 161-182.
- 550 Richards, P.C., Gatliff, R.W., Quinn, M.F., Williamson J.P. and Fannin, N.G.T. 1996b. The  
551 geological evolution of the Falkland Islands continental shelf. *From Storey, B.C., King, E.C.*  
552 *and Livermore, R.A. (eds), 1996, Weddell Sea Tectonics and Gondwana Break-up,*  
553 *Geological Society Special Publication No 108*, 105-128.
- 554 Richardson, N.J. and Underhill, J.R., 2002. Controls on the structural architecture and  
555 sedimentary character of syn-rift sequences, North Falkland Basin, South Atlantic. *Marine*  
556 *and Petroleum Geology*, *19(4)*, pp.417-443.
- 557 Sandwell, D.T., Müller, R.D., Smith, W.H.F., Garcia, E. and Francis, R. 2014. New global  
558 marine gravity model from CryoSat-2 and Jason-1 reveals buried tectonic structure. *Science*,  
559 **346** (6205), 65-67
- 560 Stanca, R.M., Paton, D.A., Hodgson, D.M., McCarthy, D.J. and Mortimer, E.J., 2019. A  
561 revised position for the rotated Falkland Islands microplate. *Journal of the Geological*  
562 *Society*, pp.jgs2018-163.
- 563 Stone, P., Richards, P.C., Kimbell, G.S., Esser, R.P. and Reeves, D., 2008. Cretaceous  
564 dykes discovered in the Falkland Islands: implications for regional tectonics in the South  
565 Atlantic. *Journal of the Geological Society*, *165(1)*, pp.1-4.
- 566 Storey, B.C., Curtis, M.L., Ferris, J.K., Hunter, M.A. and Livermore, R.A., 1999.  
567 Reconstruction and break-out model for the Falkland Islands within Gondwana. *Journal of*  
568 *African Earth Sciences*, *29(1)*, pp.153-163.
- 569 Thomson, K., 1998. When did the Falklands rotate?. *Marine and Petroleum Geology*, *15(8)*,  
570 pp.723-736.
- 571 Thomson, K. and Underhill, J. 1999. Frontier exploration in the South Atlantic: structural  
572 prospectivity in the North Falkland Basin. *AAPG Bulletin*, **83(5)**, 778-797.

573 Williams, L.S. 2015. Sedimentology of the Lower Cretaceous reservoirs of the Sea Lion  
574 Field, North Falkland Basin. *Petroleum Geoscience*, **21**, 183-198.

575 Williams, L.S. and Newbould, R., 2015. Introduction to the North Falkland Basin revisited:  
576 exploration and appraisal of the Sea Lion Field. *Petroleum Geoscience*, *21*(2-3), pp.83-84.

577 Vail, P.R., Mitchum Jr, R.M. and Thompson III, S., 1977. Seismic stratigraphy and global  
578 changes of sea level: Part 3. Relative changes of sea level from Coastal Onlap: section 2.

579 Application of seismic reflection Configuration to Stratigraphic Interpretation.

580 **Figure Captions** (all figures should be in colour for print)

581 Figure 1. Geological map of the offshore areas around the Falkland Islands modified after  
582 Richards et al., (1996). Red blocks show extent of the 2019 hydrocarbon exploration  
583 licences. Detailed fault interpretation of the North Falkland Basin (NFB) based on Lohr and  
584 Underhill (2015). The NFB consists of several subsidiary depocentres illustrated by the N-S  
585 orientated Mesozoic-Cenozoic basins. Inset image shows the location of the Falkland  
586 Islands with respect to South America.

587 Figure 2. The structural framework underlain by Bouguer gravity data of the North Falkland  
588 Basin. The Northern sector of the North Falkland Basin (NNFB) study area is illustrated with  
589 a dashed box. Grabens include: WGS - Western Graben Splay; WH - Western High; EG -  
590 Eastern Graben; EGS - Eastern Graben Splay; IBH - Intra-Basin High; PG - Phyllis Graben.  
591 The Bouguer gravity anomaly map with wavelengths <100 km modified from McCarthy et al.,  
592 2017, with data from Sandwell et al., (2014). Onshore faults are taken from Aldiss and  
593 Edwards, 1999 and offshore fault interpretation modified from Lohr and Underhill, (2015).  
594 The white lines are seismic sections which are illustrated across the NNFB (see Figs. 6 and  
595 7) and the wider NFB (see Fig.4 and 8).

596 Figure 3. (a) E-W orientated crossline from the Company merged 3D seismic dataset, across  
597 the central part of the NFB, showing a typical cross-sectional view of the Western Graben,  
598 Orca High, Eastern Graben and Eastern Flank. The asymmetrical profile of the NFB in the

599 Eastern Graben deepens towards the main bounding fault. Seismic character is poor below  
600 3.5 seconds TWT. (b) A geoseismic section showing the six reflectors interpreted in this  
601 study that define the main tectonostratigraphic units. The main bounding fault of the Western  
602 Graben, acts as a low angle detachment for antithetic faults in this part of the basin. Faults  
603 have been inferred at depth with a red dashed line and main bounding faults are labelled  
604 with fault arrows to illustrate tectonic movement. This interpretation has been adapted from  
605 Dodd et al., (2019).

606 Figure 4. Geological summary chart for the North Falkland Basin from Devonian to recent  
607 times. A comparison is made between seismic horizons of this study, with nomenclature  
608 from Richards and Hillier (2000a), along with lithology and environment interpretations from  
609 MacAulay (2015).

610 Figure 5. Orientations of 2D seismic data interpreted during this study, along with 3D seismic  
611 data used to correlate the tectonostratigraphic packages. The location of possible  
612 hydrocarbon indicators is also shown.

613 Figure 6. (a) E-W orientated 2D seismic line FALK 2000-020 across the northern sector of  
614 the NFB, showing a typical cross-sectional view in this study area of the northern most of the  
615 Eastern Graben, Eastern Flank and the Phyllis Graben. Seismic character is often lost  
616 beyond 3 - 3.5 seconds TWT. (b) A geoseismic section of FALK 2000-020 showing the six  
617 reflectors interpreted in this study that define the main tectonostratigraphic units. Here the  
618 Eastern Graben forms a symmetrical graben with a series of tilted rotational fault blocks in  
619 the syn-rift covered by a blanket of transitional to post-rift sediment. The Phyllis Graben  
620 displays an asymmetrical profile with sediment deepening to west against a bounding fault,  
621 with a series of synthetic faults developed.

622 Figure 7. (a) E-W orientated 2D seismic line FALK 2000-013 across the southern part of this  
623 NNFB study area showing the cross-sectional view of the Eastern Graben, Eastern Graben  
624 Splay, Eastern Flank of and the Phyllis Graben. Seismic amplitude character is shown down

625 until 3 seconds TWT. (b) A geoseismic section of FALK 2000-013 showing the six reflectors  
626 interpreted in this study that define the main tectonostratigraphic units. The Eastern Graben  
627 displays a symmetrical graben profile, whilst the Eastern Graben splay displays an  
628 asymmetrical geometry, deepening to the east against a bounding fault, which is separated  
629 from the Eastern Flank by a basement terrace. To the east, the syn-rift in the Eastern  
630 Graben Splay shallows and onlaps the Intra-Basin High. The Phyllis Graben displays an  
631 asymmetry deepening to the west against the boundary fault, which forms the eastern edge  
632 of the Eastern Flank. There are some discrete, high amplitude features in the late syn-rift,  
633 which could represent potential stratigraphic traps for hydrocarbons.

634 Figure 8. A geoseismic section of N-S composite line using the 2D seismic from this study  
635 and the 3D company merged 3D seismic volume along strike of the Eastern Graben of the  
636 North Falkland Basin. This line shows the six reflectors that define the main  
637 tectonostratigraphic units. In addition, high amplitude packages compare the Sea Lion  
638 discovery with the geometries observed in the Northern Lead, both within the Early Post-Rift  
639 interval. It shows a general deepening of syn-rift sediment towards to the south directly  
640 beneath Sea Lion, whilst the transitional and post-rift sediment seems to be a consistent  
641 depth across the basin. In the early post-rift, the most prospective interval, the Northern  
642 Lead, represents high amplitude package directly beneath southerly prograding clinoforms,  
643 which is similar to that of Sea Lion.

644 Figure 9. A series of two-way-travel time (milliseconds) structural interpretation maps  
645 showing different tectonostratigraphic reflectors at various stages of basin evolution. Note  
646 the black polygons indicate the fault trends, whilst the black dashed lines represent sub-  
647 cropping of the unit. Areas of white demonstrate no sediment deposition at this time. (a) Top  
648 Basement (b) Top Early Syn-Rift (c) Top Late Syn-Rift (d) Transitional/Sag Phase (e) Top  
649 Early Post-Rift and (d) Top Late Post-Rift



650 Figure 10. Seismic anomalies found on the 2D FALK 2000 seismic dataset in the study area.

651 (a) Bright spot amplitudes and potential fault trap (b) Gas chimneys escaping towards the  
652 surface, with visible paleo-pockmarks (c) Bright discontinuous amplitude anomalies that  
653 could represent hydrocarbon filled stratigraphic traps.

654 Figure 11. Schematic cross-section showing the proposed petroleum systems in the NNFB.

655 The main source rock is located in organic-rich lacustrine deposits found within the central  
656 depocentre of each graben. There is a risk of hydrocarbon leakage along major basement  
657 bounding faults within the post-rift section. The Phyllis Graben represents a potentially new  
658 area for exploration, with analogous geometries, scale and sedimentary fill to the Eastern  
659 Graben. With additional seismic acquisition, this area could represent significant future  
660 exploration in the North Falkland Basin.

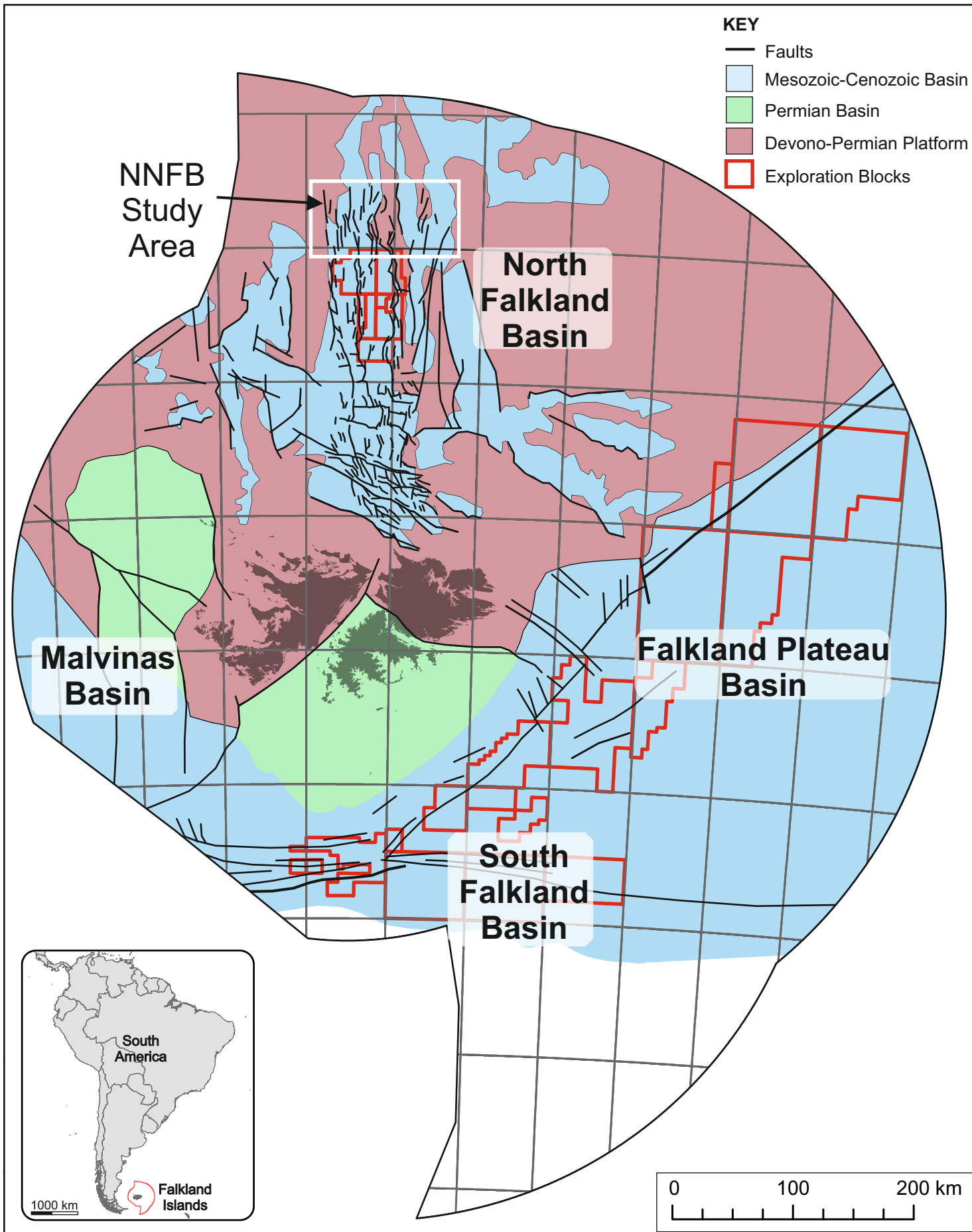


Figure. 1

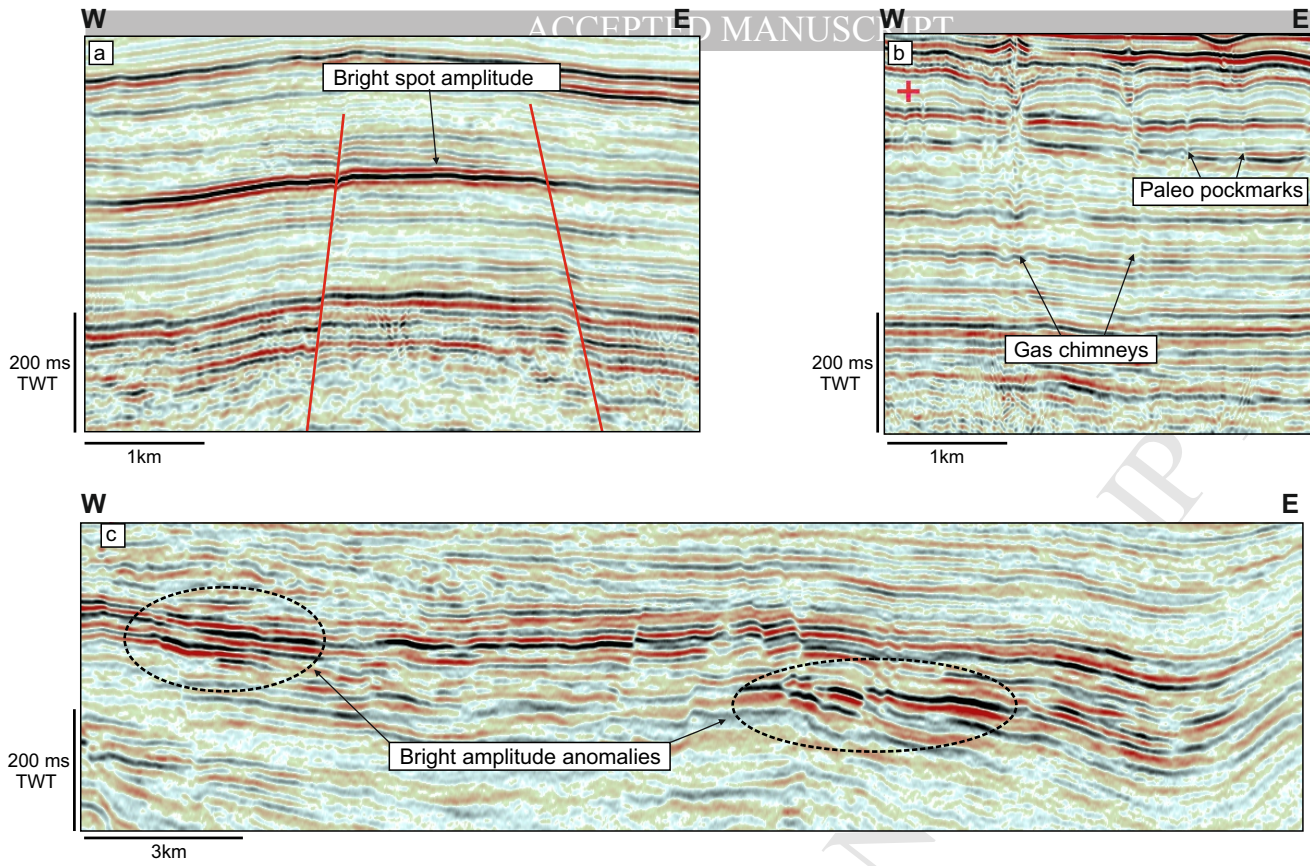


Figure. 10

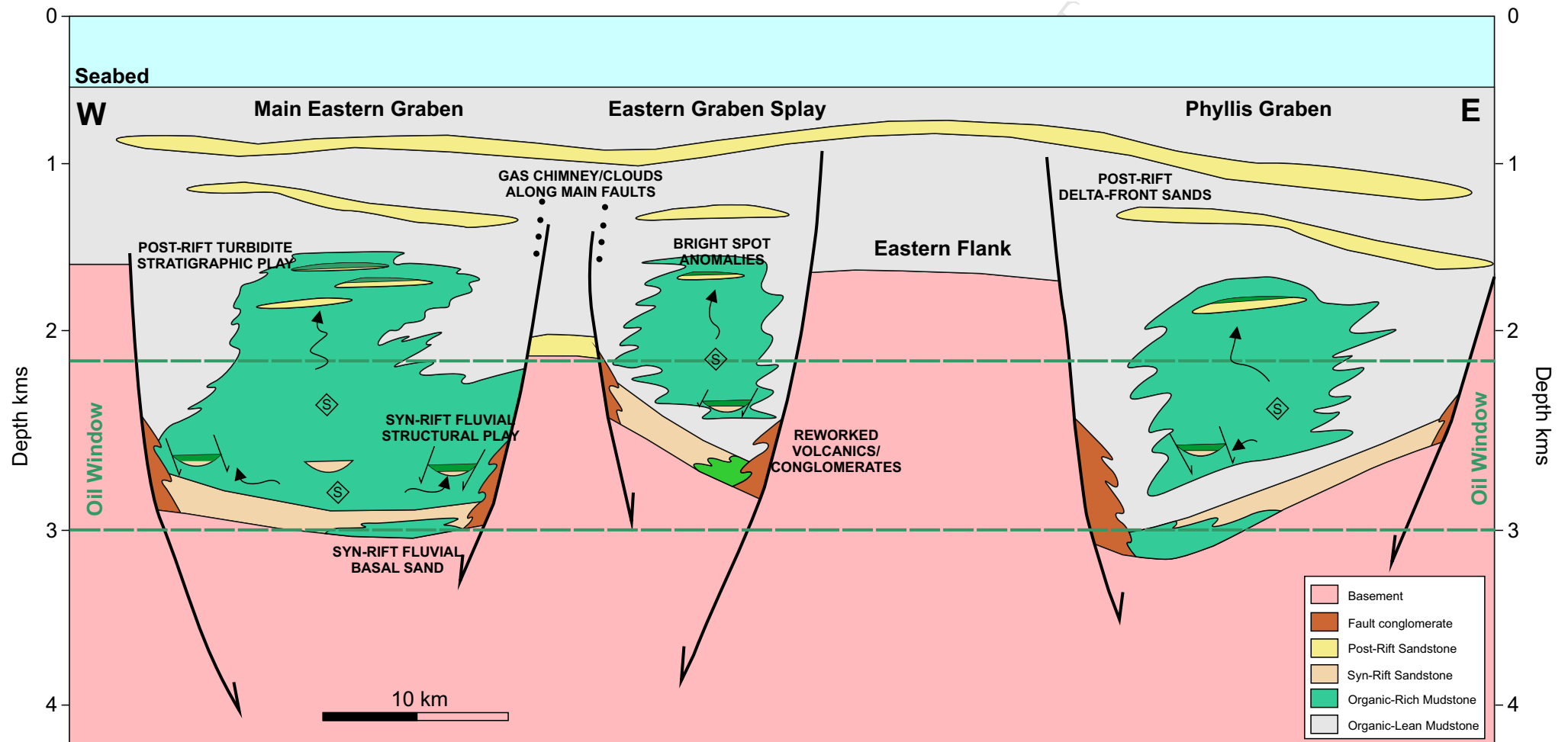


Figure. 11



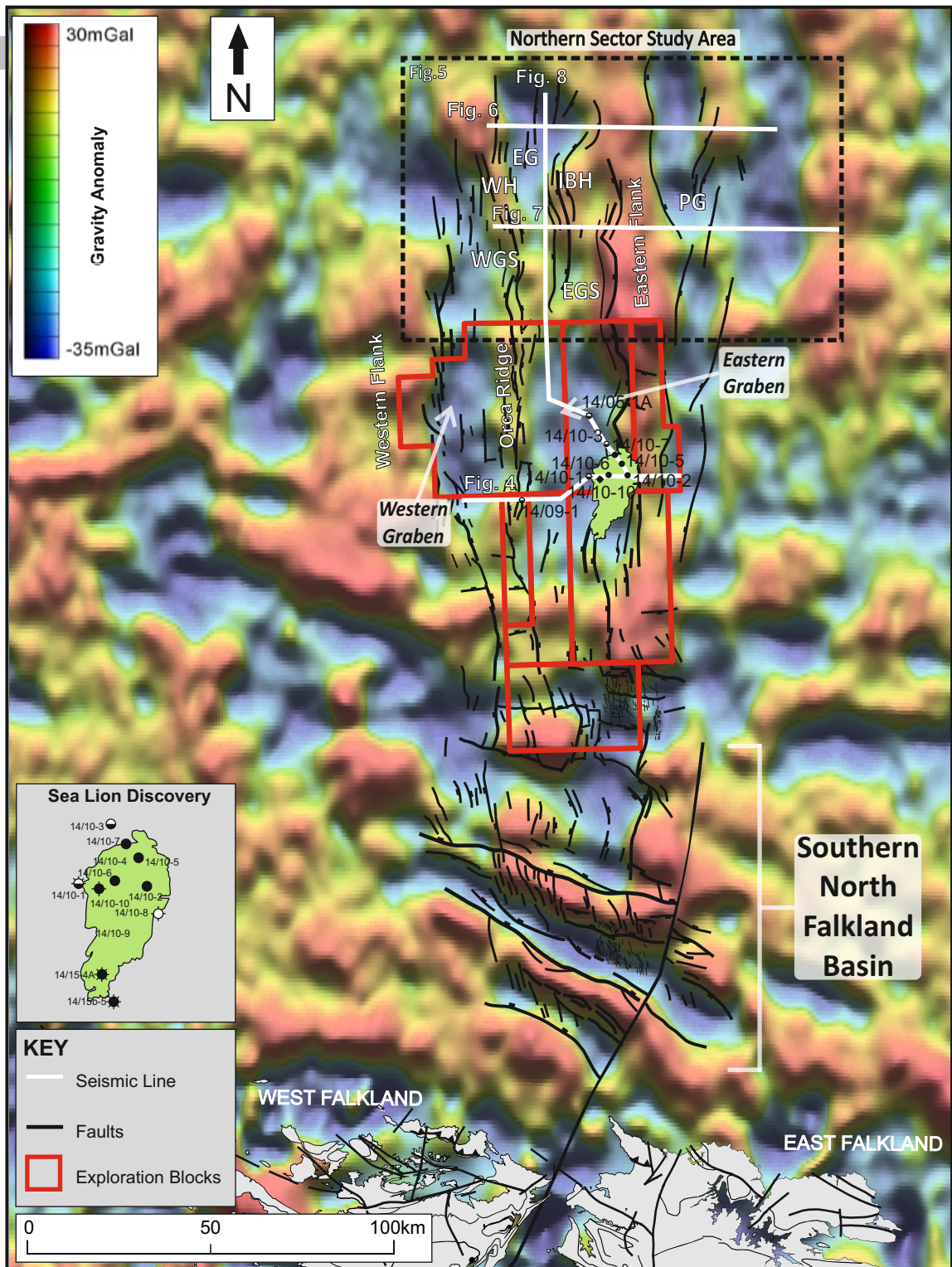


Figure. 2



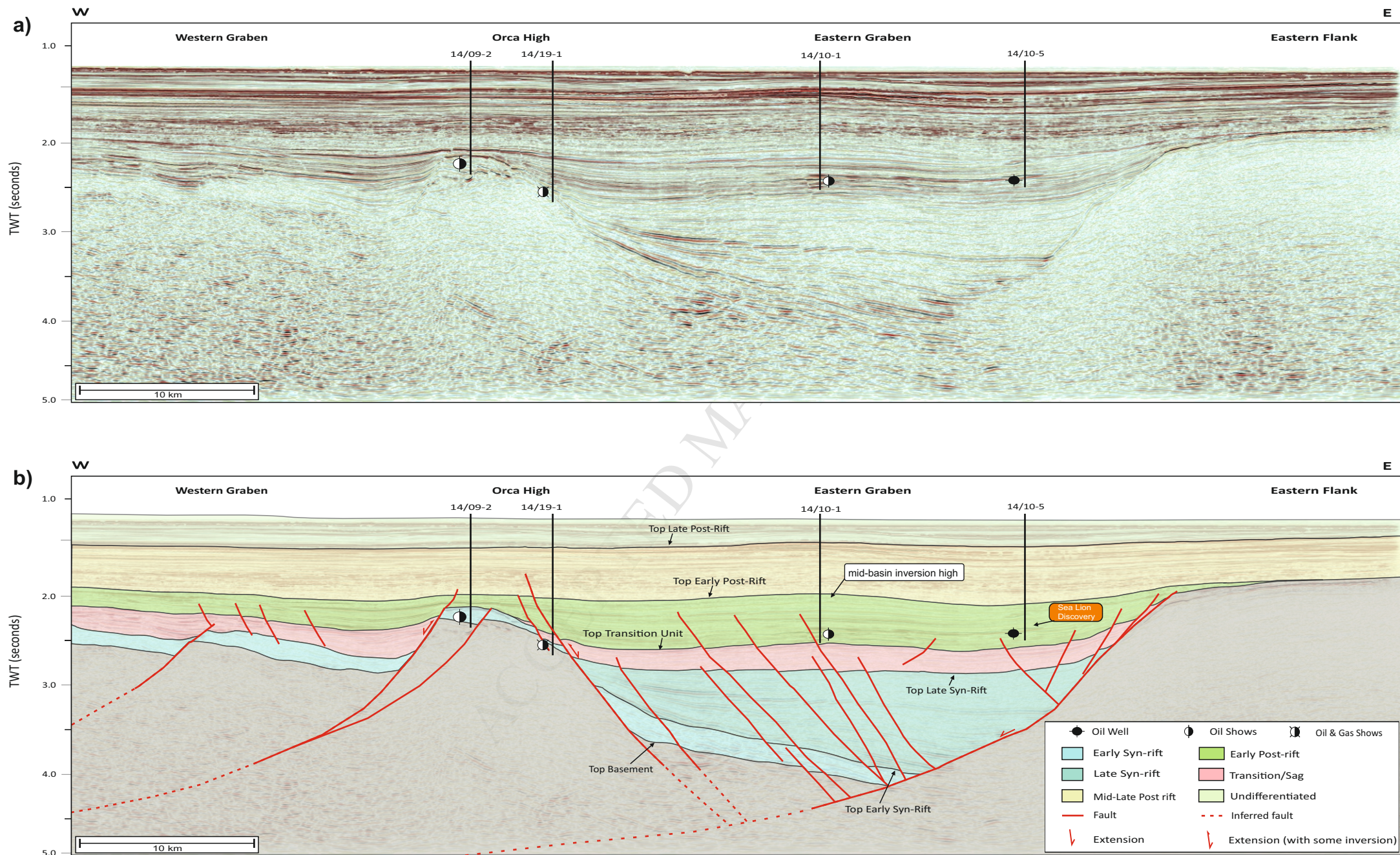


Figure. 3

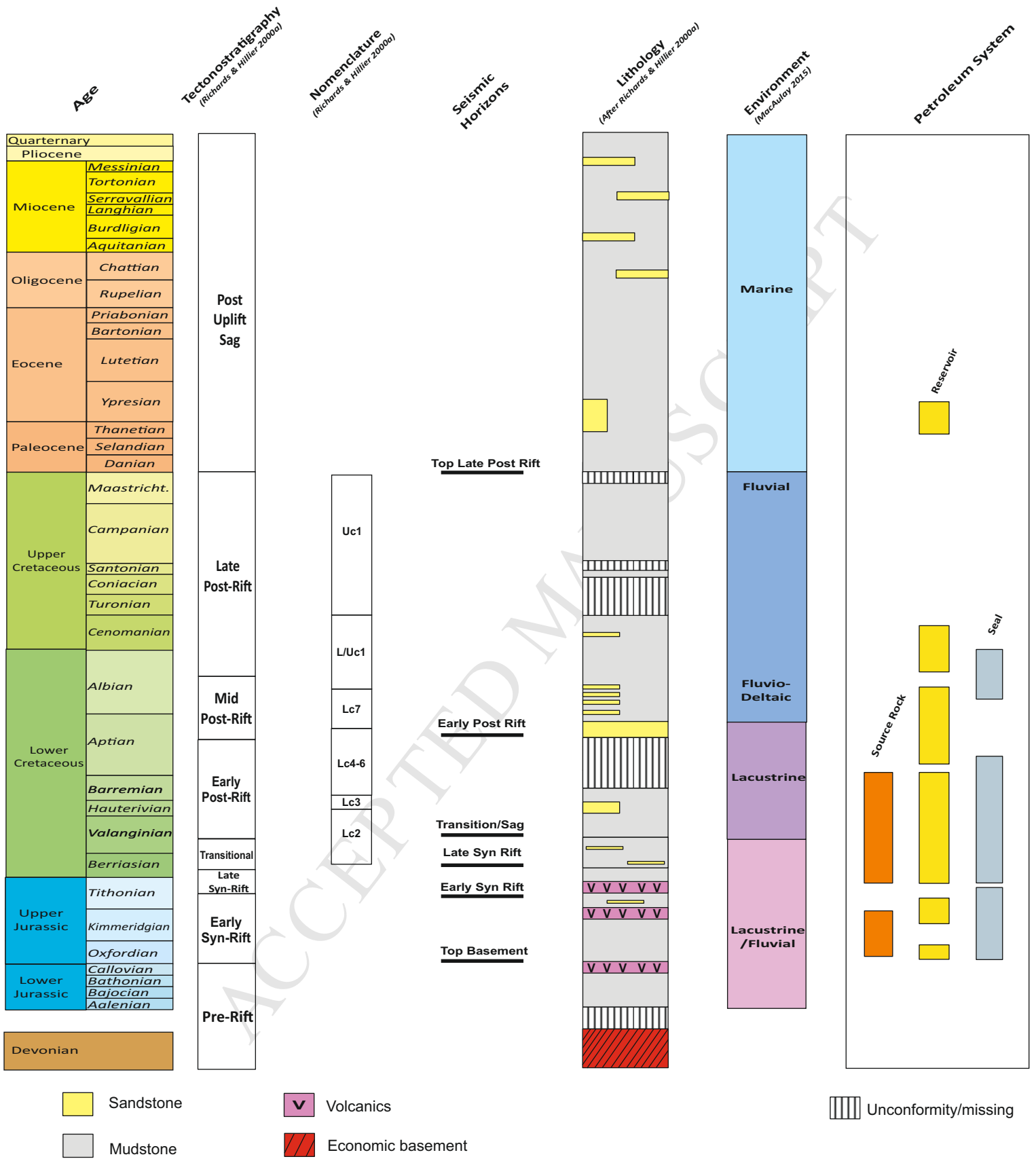


Figure. 4



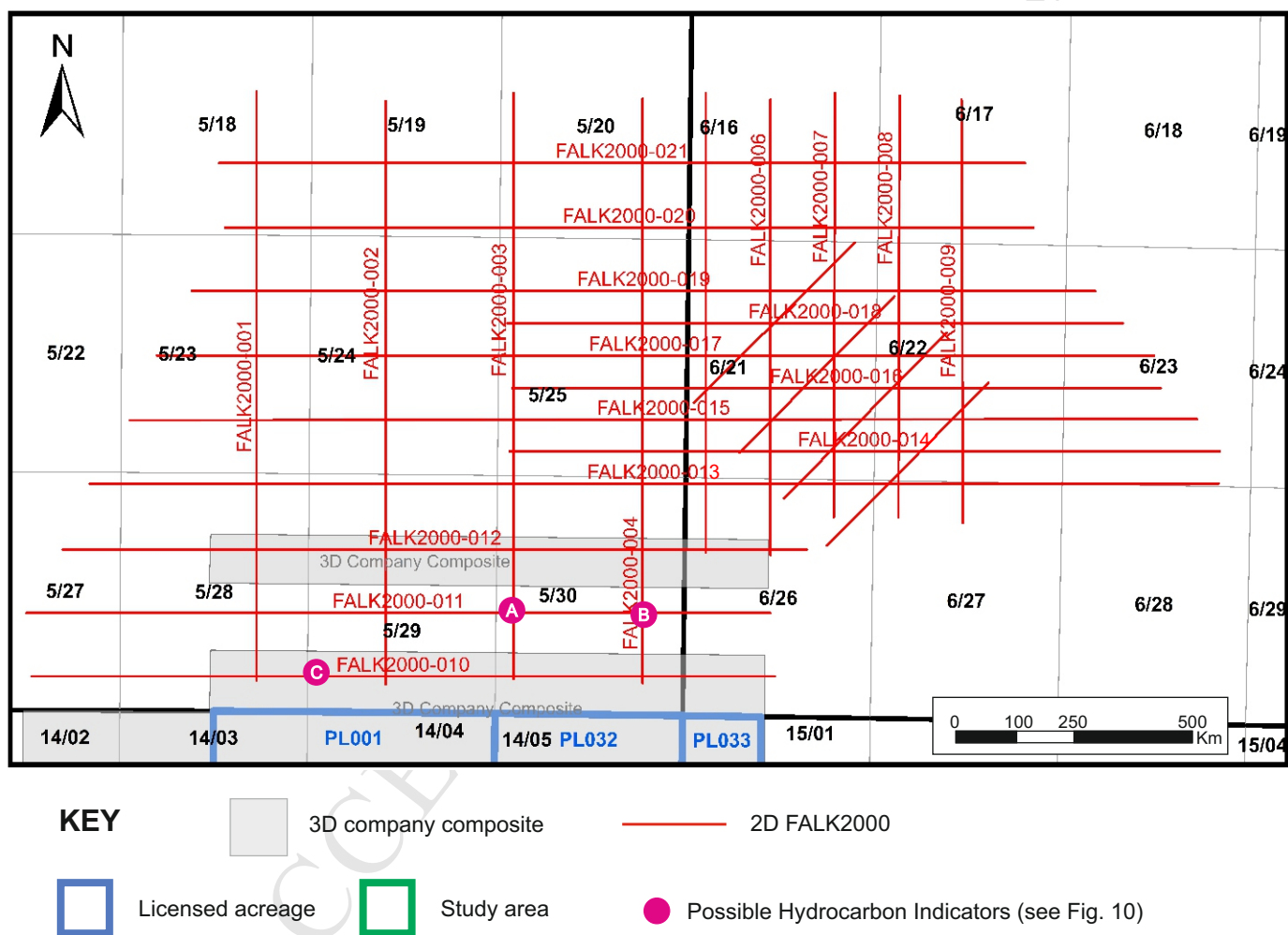


Figure. 5



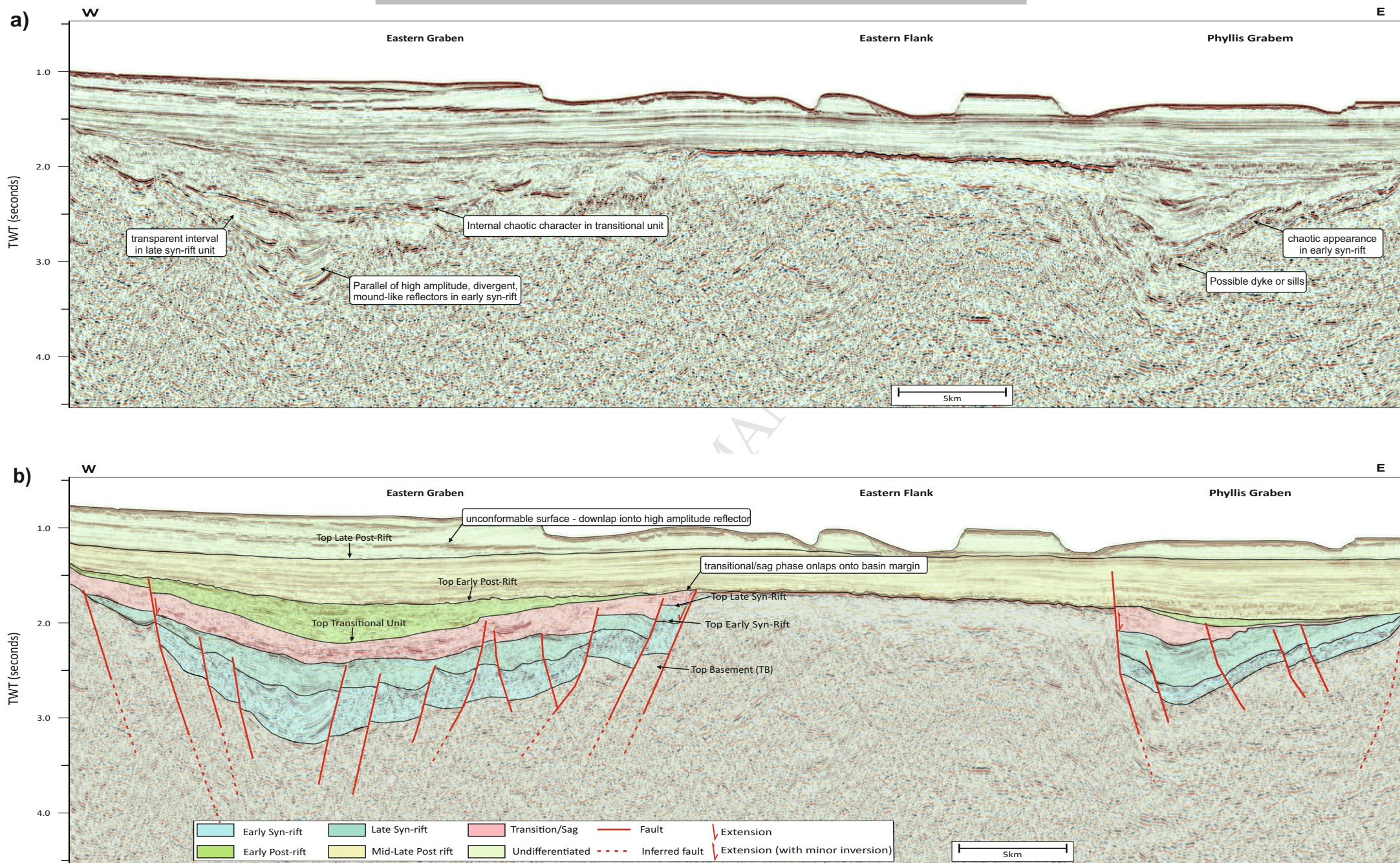


Figure. 6



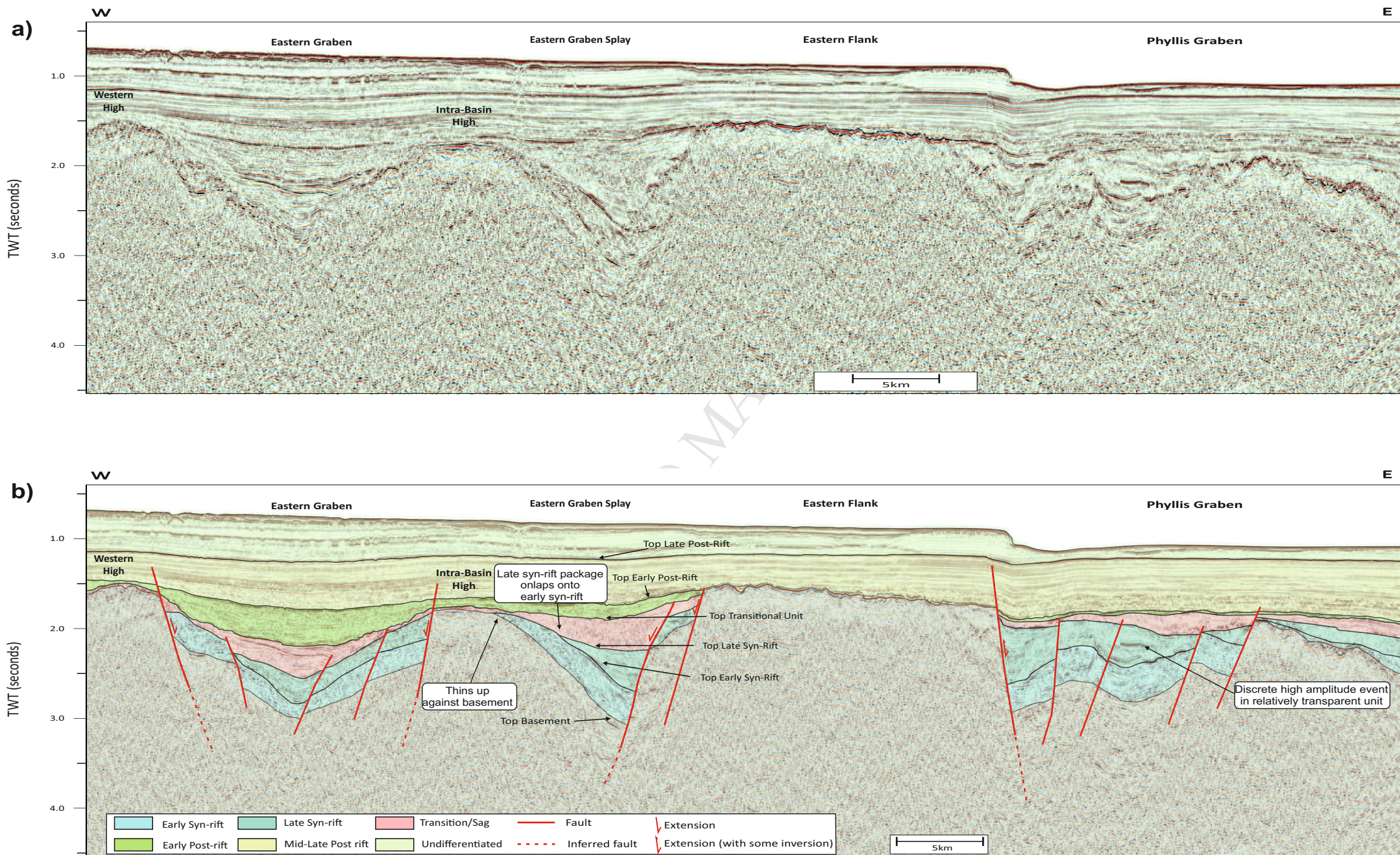


Figure. 7



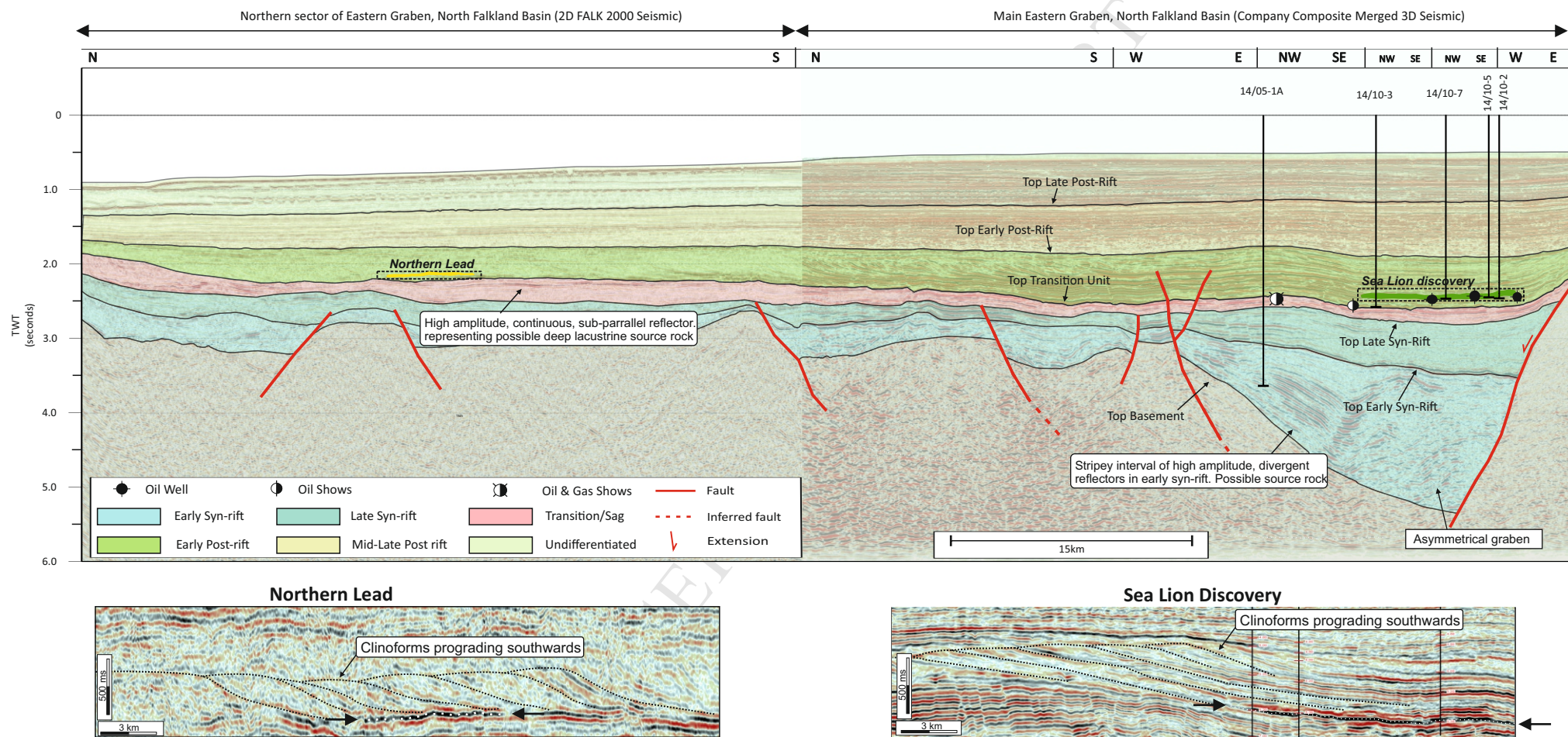


Figure. 8

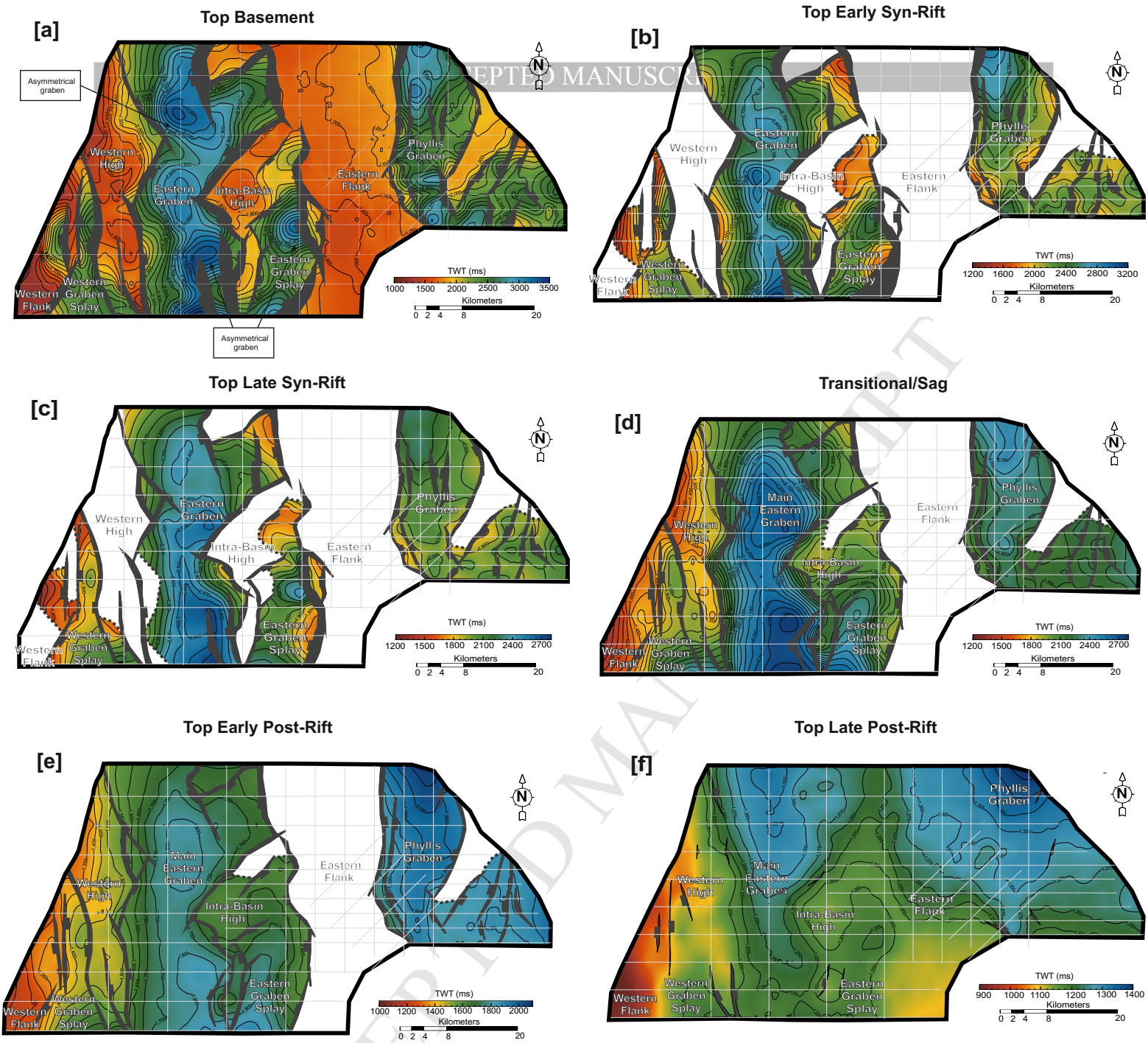


Figure. 9

## Highlights

- Four main depocentres in the NNFB are identified in 2D seismic data
- NW-SE and NE-SW trending faults controlled syn-rift graben development
- Six tectonostratigraphic units are identified and described across the NNFB
- Identification of hydrocarbon plays within early post-rift and syn-rift
- The NNFB forms an extension of the proven NFB petroleum system